

1-14-2015

# Full vector low-temperature magnetic measurements of geologic materials

Joshua M. Feinberg

*University of Minnesota - Twin Cities, feinberg@umn.edu*

Peter A. Solheid

*University of Minnesota - Twin Cities*

Nicholas L. Swanson-Hysell

*University of Minnesota - Twin Cities*

Mike J. Jackson

*University of Minnesota - Twin Cities, jacks057@umn.edu*

Julie A. Bowles

*University of Wisconsin-Milwaukee, bowlesj@uwm.edu*

Follow this and additional works at: [https://dc.uwm.edu/geosci\\_facart](https://dc.uwm.edu/geosci_facart)



Part of the [Earth Sciences Commons](#)

---

## Recommended Citation

Feinberg, Joshua M.; Solheid, Peter A.; Swanson-Hysell, Nicholas L.; Jackson, Mike J.; and Bowles, Julie A., "Full vector low-temperature magnetic measurements of geologic materials" (2015). *Geosciences Faculty Articles*. 11.  
[https://dc.uwm.edu/geosci\\_facart/11](https://dc.uwm.edu/geosci_facart/11)

This Article is brought to you for free and open access by UWM Digital Commons. It has been accepted for inclusion in Geosciences Faculty Articles by an authorized administrator of UWM Digital Commons. For more information, please contact [open-access@uwm.edu](mailto:open-access@uwm.edu).



# Geochemistry, Geophysics, Geosystems

## RESEARCH ARTICLE

10.1002/2014GC005591

### Special Section:

Magnetism From Atomic to Planetary Scales: Physical Principles and Interdisciplinary Applications in Geo- and Planetary Sciences

### Key Points:

- New low-temperature magnetic instrument
- Full vector records across magnetic transitions
- Applications to paleomagnetic studies

### Supporting Information:

- Readme
- Supporting data and code

### Correspondence to:

J. M. Feinberg,  
feinberg@umn.edu

### Citation:

Feinberg, J. M., P. A. Solheid, N. L. Swanson-Hysell, M. J. Jackson, and J. A. Bowles (2015), Full vector low-temperature magnetic measurements of geologic materials, *Geochem. Geophys. Geosyst.*, 16, doi:10.1002/2014GC005591.

Received 26 SEP 2014

Accepted 6 JAN 2015

Accepted article online 14 JAN 2015

## Full vector low-temperature magnetic measurements of geologic materials

Joshua M. Feinberg<sup>1</sup>, Peter A. Solheid<sup>1</sup>, Nicholas L. Swanson-Hysell<sup>1,2</sup>, Mike J. Jackson<sup>1</sup>, and Julie A. Bowles<sup>1,3</sup>

<sup>1</sup>Institute for Rock Magnetism, Department of Earth Sciences, University of Minnesota, Minneapolis, Minnesota, USA,

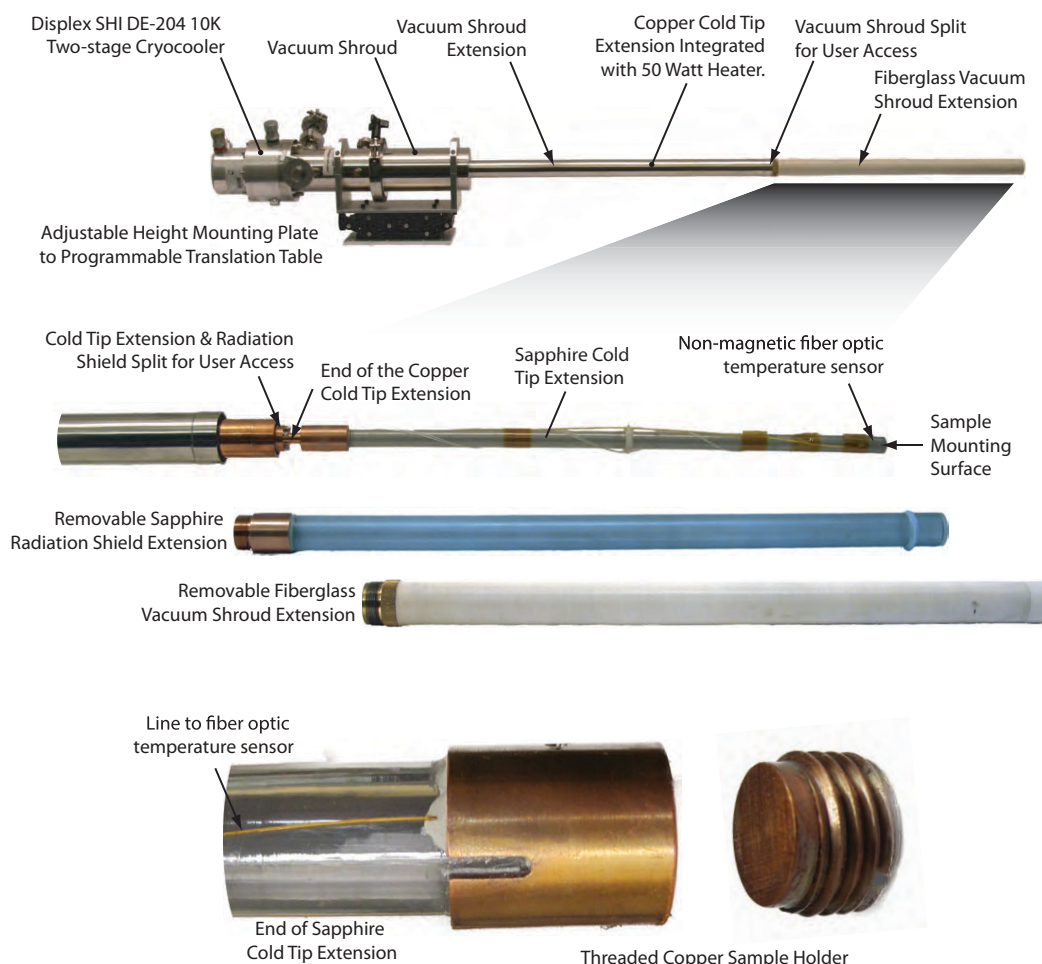
<sup>2</sup>Department of Earth and Planetary Science, University of California, Berkeley, California, USA, <sup>3</sup>Department of Geosciences, University of Wisconsin-Milwaukee, Milwaukee, Wisconsin, USA

**Abstract** The magnetic properties of geologic materials offer insights into an enormous range of important geophysical phenomena ranging from inner core dynamics to paleoclimate. Often it is the low-temperature behavior (<300 K) of magnetic minerals that provides the most useful and highest sensitivity information for a given problem. Conventional measurements of low-temperature remanence are typically conducted on instruments that are limited to measuring one single-axis component of the magnetization vector and are optimized for measurements in strong fields. These instrumental limitations have prevented fully optimized applications and have motivated the development of a low-temperature probe that can be used for low-temperature remanence measurements between 17 and 300 K along three orthogonal axes using a standard 2G Enterprises SQUID rock magnetometer. In this contribution, we describe the design and implementation of this instrument and present data from five case studies that demonstrate the probe's considerable potential for future research: a polycrystalline hematite sample, a polycrystalline hematite and magnetite mixture, a single crystal of magnetite, a single crystal of pyrrhotite, and samples of Umkondo Large Igneous Province diabase sills.

## 1. Introduction

Magnetic behavior at low temperatures (<300 K) is one of the most sensitive indicators of the iron mineral phases and their concentrations and grain size distributions in natural samples. Changes in magnetocrystalline anisotropy and crystallographic structure give rise to low-temperature transitions that are diagnostic of specific mineral phases. The Morin transition of hematite (at ~262 K) [Morin, 1950], the Verwey transition of magnetite (at ~122 K) [Verwey, 1939], and the Besnus transition of pyrrhotite (at ~32 K) [Besnus and Meyer, 1964] are all diagnostic of common magnetic minerals that carry remanence at Earth surface temperatures. Other phases that acquire remanence at low temperature, such as siderite (with a Néel temperature of 38 K) [Frederichs et al., 2003] and superparamagnetic grains [Worm and Jackson, 1999], can also be readily identified through their low-temperature behavior. In addition to the utility of low-temperature data as a diagnostic tool for magnetic mineral identification and characterization, irreversible changes in remanence that are associated with cycling to low temperatures are often used as a tool in paleomagnetic studies. Low-temperature steps in paleodirectional and paleointensity study are applied in some protocols with the goal of preferentially removing magnetic remanence held by multidomain grains and thereby isolating magnetizations held by single-domain grains [e.g., Schmidt, 1993; Dunlop, 2003; Yamamoto et al., 2003].

Low-temperature remanence experiments are routinely conducted on the Quantum Designs Magnetic Properties Measurement System (MPMS), most often with the intention of revealing information about the dominant magnetic mineral phases and grain size distribution. While these instruments are adept at a range of low-temperature experiments, understanding the full behavior of a rock's natural remanence at low temperature is hampered by the measurement capabilities being limited to a single axis and the instrument not providing an ultralow field environment. If low-temperature measurements of a natural remanence (NRM) are desired using such instrumentation, great care must be taken to align the NRM with the measurement axis, and any directional change during thermal cycling will not be captured. In such an instrument, deviation from the single axis will result in a measured magnetization that is less than the specimen's actual total magnetization.



**Figure 1.** Components of the Institute for Rock Magnetism's low-temperature instrument (IRM-LTI). The length of the instrument is 175 cm from the back of the cryocooler to the tip of the fiberglass vacuum shroud.

In this contribution, we describe a low-temperature probe developed for use with superconducting rock magnetometers (SRM) at the Institute for Rock Magnetism (IRM), University of Minnesota. This instrument allows for three-axis full-vector measurements of magnetic remanence at temperatures between 300 and 17 K in low-field environments ( $<10$  nT). It was developed with different engineering, but the same intent, as a low-temperature insert that has previously been implemented at the University of Rochester Paleomagnetic Laboratory [Smirnov and Tarduno, 2011].

## 2. Instrument Design

In order to develop the capacity to make three-axis measurements of remanence in an ultralow-field environment, a cryostat insert was developed at the IRM in cooperation with ColdEdge Technologies (Allentown, PA) (Figure 1). This low-temperature instrument (IRM-LTI) allows for three-axis measurements to be made between room temperature and  $\sim 17$  K using horizontal-loading SRMs. There are many advantages to outfitting a superconducting rock magnetometer for measurements at low temperatures. First, these instruments are specifically designed for three-axis remanence measurements, and ambient fields are minimized using a superconducting lead shield. Nulling fields are applied by external coils while the shield cools to superconducting temperatures, ultimately trapping a  $\sim 2$ – $3$  nT field along all three cardinal directions. Moreover, modern SRM instruments utilize DC-SQUID sensors that offer greater sensitivity than existing low-temperature magnetometers and susceptometers that rely on RF-SQUID sensors. A room temperature, open-ended bore is present in all 2G Enterprises magnetometers, allowing samples to be easily moved into and out of the measurement region of the instrument.

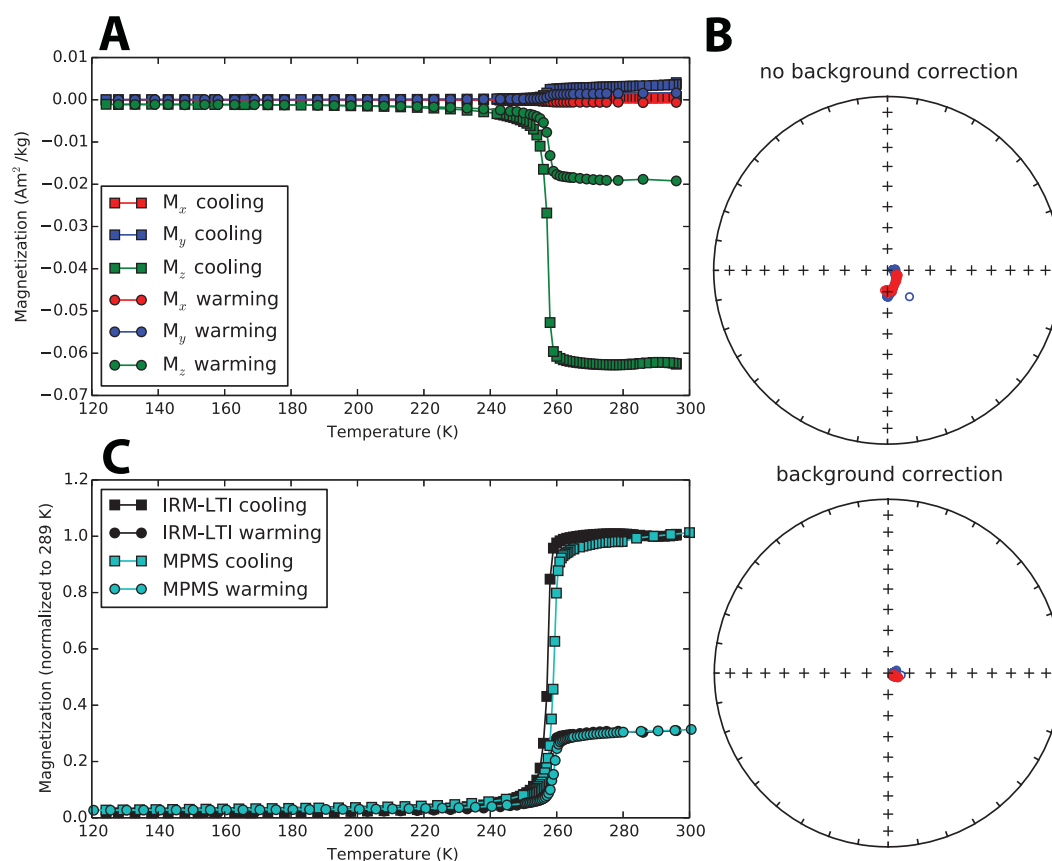
The cryogenic insert is cooled by a pneumatically driven SHI Displex SH-204 10 K two-stage cryocooler with a 73.7 cm copper cold tip extension and 3.8 cm diameter stainless steel vacuum shroud (Figure 1). A 40.6 cm sapphire cold tip extension and sapphire radiation shield are joined to the end of the copper extension in order to prevent any radio frequency noise from traveling down the copper extension tip into the measurement region of the magnetometer. Sapphire at low temperatures has a high thermal conductivity, comparable to that of copper below 100 K. A fiberglass vacuum shroud extension mounts to the stainless steel shroud to provide a nonmetallic extension to the vacuum insulation space. Additional temperature control is provided by a noninductively wound, 50 W cartridge heater integrated into the body of the probe and mounted out of the sensor region, 40.6 cm away from the sample. The temperature inside the probe is monitored directly at the sample using a nonmagnetic, fiber-optic temperature sensor (NeoOptics T1 probe) and two silicon diode sensors, one mounted on the copper cold tip extension near the heater and one mounted on the first stage of the cryocooler to monitor the radiation shield temperature. The probe assembly is mounted on a computer-controlled translation table driven by a stepper motor for automated movement in and out of the magnetometer, enabling background and sample measurements to be made throughout cooling/warming cycles.

Samples can be affixed to the end of the sapphire extension tip in a variety of ways. The goal is to maximize the thermal contact conductance between a sample and the sapphire extension tip. In its simplest form, samples can be affixed using Kapton tape (a polyimide film with silicone bonding agent that maintains its adhesive properties even at cryogenic temperatures), but at temperatures below  $\sim 50$  K, samples may lose efficient thermal contact with the sapphire extension tip. Alternatively, a threaded copper sample holder can be bonded to the end of the sapphire extension tip, allowing samples to be held in thermal contact with the sapphire (Figure 1). Both of these methods allow for samples with maximum cylindrical dimensions of 9 mm diameter and 5 mm height.

Measurements without any sample show maximum magnetizations of  $\sim 3 \times 10^{-8}$  A m<sup>2</sup> with some degree of temperature dependence (see supporting information). The temperature-dependent magnetization of the IRM-LTI is largely repeatable, and empty probe experiments are conducted prior to measurement runs to enable background subtraction if deemed necessary. Samples can be cooled from room temperature to  $\sim 17$  K over the course of about 3 h and then warmed back to room temperature over another 3–4 h. Thus, the time required to collect low-temperature data using the IRM-LTI is roughly 3 times that required for the simplest low-temperature experiments on the MPMS. Alternatively, a user can define a more narrow temperature range for more detailed studies of low-temperature phenomena.

The approach and technical design of the LTI in this study (IRM-LTI) builds on ideas of an earlier effort by the late William Goree of 2G Enterprises along with Aleksey Smirnov and John Tarduno [Smirnov and Tarduno, 2011]. The key innovation of that collaborative effort was the design of a low-temperature instrument (that we will refer to as the ST-LTI) that could be inserted into a standard bore 2G Enterprises cryogenic rock magnetometer. In the ST-LTI, the sample was cooled by direct conduction in a bath of continually flowing liquid He. The sensors of the magnetometer were protected from the cooled low-temperature probe by a double-wall fiberglass tube, which was continually evacuated by a vacuum pump during measurements. Sample temperatures were monitored by a diode temperature sensor, which by necessity, was located about 10 cm behind the sample in order to avoid interference with the magnetometer pickup coils.

The design of the LTI in this study (IRM-LTI) has several advantages and disadvantages when compared to the ST-LTI. One comparative disadvantage of the IRM-LTI is that the volume of instrumental material introduced into the measurement region of the magnetometer is greater than that of the ST-LTI, and background signals are consequently larger. The IRM-LTI is also intrinsically somewhat slower than the ST-LTI or the MPMS, since heat must be transported by solid-state conduction through a long rod rather than by advection in a fluid. Comparative advantages include: (1) because the IRM-LTI is moved via a translation table during measurements, it is possible to quantify and adjust for background magnetization and magnetometer drift for each measurement; (2) this translation allows for the possibility of in-line treatments, such as alternating-field demagnetization and acquisition of anhysteretic remanent magnetization; and (3) utilizing a closed-cycle cryocooler systems allows the IRM-LTI to operate free of liquid helium. An additional positive aspect of the IRM-LTI is that it is currently available for community use through the Institute for Rock Magnetism's Visiting Fellowship Program (<http://www.irm.umn.edu/IRM/applying.html>).



**Figure 2.** Low-temperature thermal cycling of IRMs for a sample of polycrystalline Labrador hematite. (a) Background-corrected three-axis remanence data for a 1.2 T IRM cycled to and from 120 K using the IRM-LTI. There is prominent remanence loss as the specimen cooled through the Morin transition ( $\sim 258$  K). (b) Equal-area plots of the uncorrected and corrected directions throughout thermal cycling. Blue points are during cooling while red points are during warming. All data plot in the upper hemisphere. (c) Comparison between the total vector magnetization calculated from the three-axis IRM-LTI and one-axis data collected for a similar experiment conducted on an MPMS.

### 3. Case Studies Using the Instrument

The five case studies that follow serve three primary purposes:

1. To demonstrate that the temperature-dependent intensity data acquired using the IRM-LTI is virtually identical to that acquired using MPMS instruments (as shown by *Smirnov and Tarduno* [2011] for the ST-LTI), which should give confidence to potential users.
2. To demonstrate the utility of three-axis full-vector magnetic measurements in several different rock magnetic and paleomagnetic applications. In the course of doing so, novel results have been obtained.
3. To inspire new uses for the IRM-LTI, and in some instances, we have provided suggestions for future studies that we have not yet explored.

#### 3.1. Polycrystalline Hematite

A metamorphic schist containing polycrystalline hematite from Labrador, Canada with a strong foliation was pulsed in a 1.2 T field approximately perpendicular to the foliation plane to impart an isothermal remanent magnetization (IRM) (using an ASC Model IM-10–30 Impulse Magnetizer). The strong shape-preferred alignment of the hematite grains in this polycrystalline sample allows us to explore whether its remanence direction shifts after cooling through the Morin transition. Room temperature remanence in a single crystal of hematite falls within the basal plane, while remanence below the Morin transition is oriented along the normal to this plane. After being imparted, the IRM was cycled from room temperature to  $\sim 130$  K using the IRM-LTI in conjunction with an SRM (Figure 2a). The specimen underwent a significant loss of remanence as it cooled through the Morin transition which is partially recovered as the sample warmed back through the

transition temperature (Figures 2a and 2c). During this low-temperature cycling, the sample's directional remanence remained subparallel to the foliation normal (i.e., on average approximately parallel to the  $c$  axes of the individual crystals) and did not migrate toward the plane of the magnetic fabric (Figure 2b). Even without any background correction, the majority of the remanence lies close to the  $z$  axis of the stereonet. However, there is a  $15^\circ$  deflection in inclination away from  $90^\circ$  (Figure 2b). This deflection can be attributed to the magnetization of the IRM-LTI itself. The IRM-LTI's magnetization and low-temperature behavior can be important when measuring weak remanence samples and background subtraction is necessary. This particular experiment was conducted when the IRM-LTI was outfitted with a relatively magnetic silicon diode thermometer that contributed to a background signal of  $\sim 10^{-6}$  A m<sup>2</sup>. Subsequent improvements to the instrumental setup have led to the current use of a far less magnetic fiber-optic temperature sensor that contributes  $\sim 10^{-8}$  A m<sup>2</sup> to the background signal. By measuring the low-temperature cycling of the IRM-LTI without a mounted sample, we can characterize the background of the instrument and then subtract this influence from the measured sample data. When this background subtraction is done for the hematite sample, the data are tightly grouped near  $90^\circ$  of inclination, illustrating the uniformity of the magnetization in the direction of the applied IRM before and after cooling through the Morin transition (Figure 2b). The fact that the remanence remains parallel to the original applied field direction throughout the experiment suggests either that the alignment of the hematite crystals was not perfect (otherwise the remanence would have fallen within the basal plane) and/or that the remanence was held by defects within hematite.

In this instance, the data collected with the three-axis IRM-LTI are very similar to those collected with a one-axis MPMS as part of a standard low-temperature cycling of room temperature saturation IRM (Figure 2c). The ratio of the room temperature remanence before and after low-temperature cycling is nearly identical. One subtle difference between the two data sets is the slightly lower estimates of the Morin transition temperature on cooling and warming from the IRM-LTI (257.4 K) as compared to the MPMS (259.4 K; see supporting information for details on these estimates of transition temperature). One possible explanation for this minor difference is a small thermal lag within the IRM-LTI. Additional measurements at a variety of cooling rates will test this explanation in the future.

### 3.2. Mixed Assemblage of Hematite and Magnetite

A polycrystalline hematite specimen containing trace amounts of magnetite was imparted with a 1.2 T IRM. Following this treatment, the specimen was cooled to  $\sim 20$  K and then warmed back to room temperature inside the probe (Figure 3). This measurement is roughly equivalent to the room temperature saturation isothermal remanent magnetization (RT-SIRM) experiments routinely conducted on MPMS instruments and allows one to explore how the remanence of each phase evolves during low-temperature cycling from room temperature to 20 K. The effects of the Morin and Verwey transitions are quite prominent as the sample was cycled to and from low temperature while the direction of the magnetization did not vary (Figure 3a).

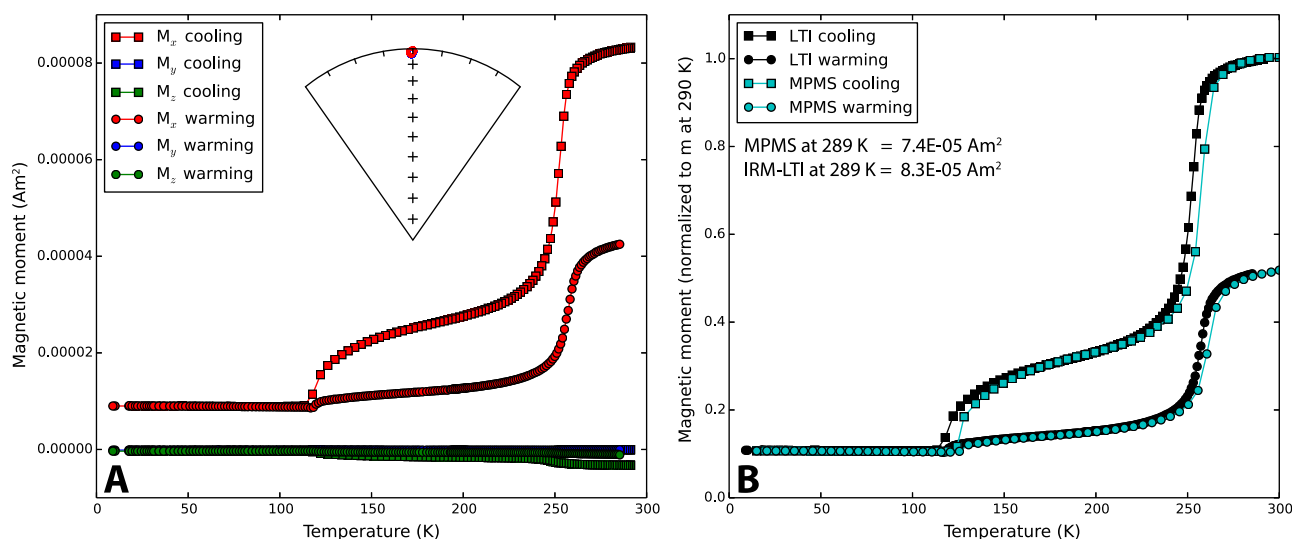
After completion of the IRM-LTI experiment, a 1.2 T IRM was again applied to the specimen and then thermally cycled using an MPMS. Comparison of the IRM-LTI and MPMS data shows the results to be quite similar (Figure 3b). Two subtle differences can be seen between the two data sets: (1) there is a systematic  $\sim 2$  K shift of the Morin and Verwey transitions toward cooler temperatures in the IRM-LTI data compared to the MPMS data (see supporting information for details), as was seen with the Labrador hematite sample (Figure 2c), and similar to the temperature offset observed by Smirnov and Tarduno in the ST-LTI experiments; (2) the magnitude of the magnetization is slightly less for the MPMS data than the total vector of the three-axis IRM-LTI data, suggesting that some of the magnetization may have been off axis during the MPMS experiment and thereby not detected by the single-axis measurement capabilities of that instrument.

In future studies, the IRM-LTI can be used to disentangle the remanence held by each of these phases by imparting specimens with multiple, mutually perpendicular magnetizations, similar in style to a Lowrie test [Lowrie, 1990], where the remanence associated with specific phases is forced to lie along different cardinal directions. In this way, the IRM-LTI may provide a unique method for nondestructive exploration of the temperature-dependent magnetic behavior of mixed magnetic mineral assemblages.

### 3.3. Single Crystal of Magnetite

A square-based pyramid was cut from a magnetite octahedron to explore the directional behavior of remanence for a single crystal of multidomain magnetite during low-temperature cycling using the IRM-LTI. The



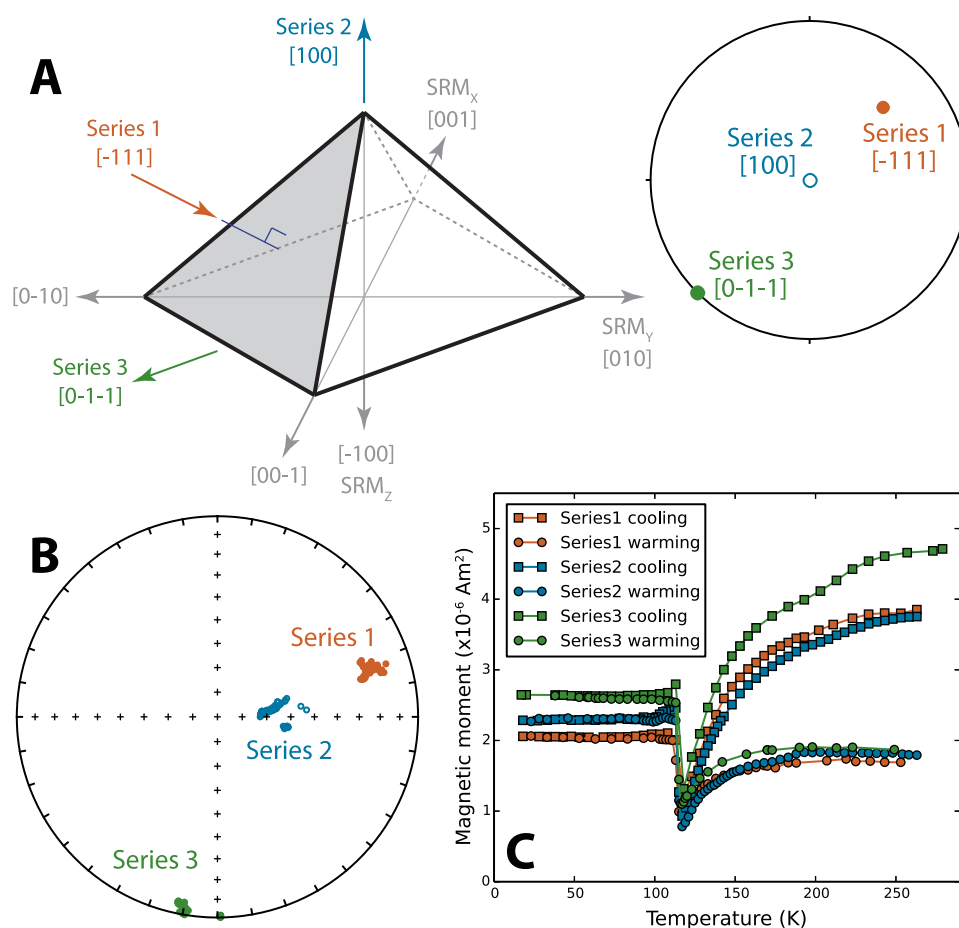


**Figure 3.** Low-temperature thermal cycling of a polycrystalline hematite sample containing trace quantities of magnetite. (a) Plot of background-corrected, three-axis remanence data for cooling to 20 K and warming back to room temperature using the IRM-LTI following application of a 1.2 T isothermal remanence. Both the Morin and Verwey transitions are clearly resolved. The inset shows a portion of an equal-area plot with the tightly grouped directions of all 114 data points measured during the thermal cycling. (b) Plot of the total magnetization measured by the IRM-LTI and single-axis data obtained on an MPMS for a 1.2 T isothermal remanence as it was cycled to and from low temperature ( $\sim 20$  K). The data are normalized to the magnetization measured at  $\sim 290$  K on each respective instrument.

2 mm subspecimen used in this study was prepared using a low-speed saw to remove the pyramid from the top of a natural magnetite octahedron that originated in the Central African Republic (mined and sold by Ikon Mining Co.). Three experiments (or “Series”) were conducted, each of which consisted of applying a strong (1.2 T) IRM in a distinct direction and then cycling that magnetization to low temperature (20 K) and back. A schematic drawing of the specimen and its orientation relative to the measurement axes of the superconducting rock magnetometer (SRM) and to the orientations of three separate 1.2 T IRMs is shown in Figure 4a. The low-temperature cycling of the isothermal remanence applied to the specimen is described in detail below, but it is important to note that there are multiple sources of orientation error that are inherent to the measurements. First, despite the care taken during specimen preparation, it is likely that the square base of the pyramid was not cut exactly parallel to the (100) of magnetite and may be misaligned by  $<10^\circ$ . Such an error will affect all of the directional IRM experiments. Second, for each IRM there may be a small  $<10^\circ$  error between the orientation of the applied field and its targeted crystallographic direction. Third, there is also a small ( $<5^\circ$ ) error associated with placement of the specimen back on the end of the IRM-LTI after each successive IRM. This latter error should only be rotational around the z axis of the magnetometer.

We attempted to impart the IRMs along the  $[-111]$ ,  $[100]$ , and  $[0-1-1]$  of magnetite because these directions correspond to magnetocrystalline easy, hard, and intermediate axes, respectively, at room temperature. As magnetite cools below the Verwey transition ( $T_V$ ), the mineral experiences a first-order phase transition where its crystal symmetry changes from cubic (c) to monoclinic (m). The orientations of the  $a$ ,  $b$ , and  $c$  axes of the resulting monoclinic magnetite will depend on the magnetic environment in which this transition occurs, but by convention are described as  $[001]_m/[001]_c$ ,  $[100]_m/[110]_c$ , and  $[010]_m/[-110]_c$  [e.g., Kasama *et al.*, 2010]. The easy, intermediate, and hard magnetocrystalline axes for monoclinic magnetite are  $[001]_m$ ,  $[010]_m$ , and  $[100]_m$ , respectively.

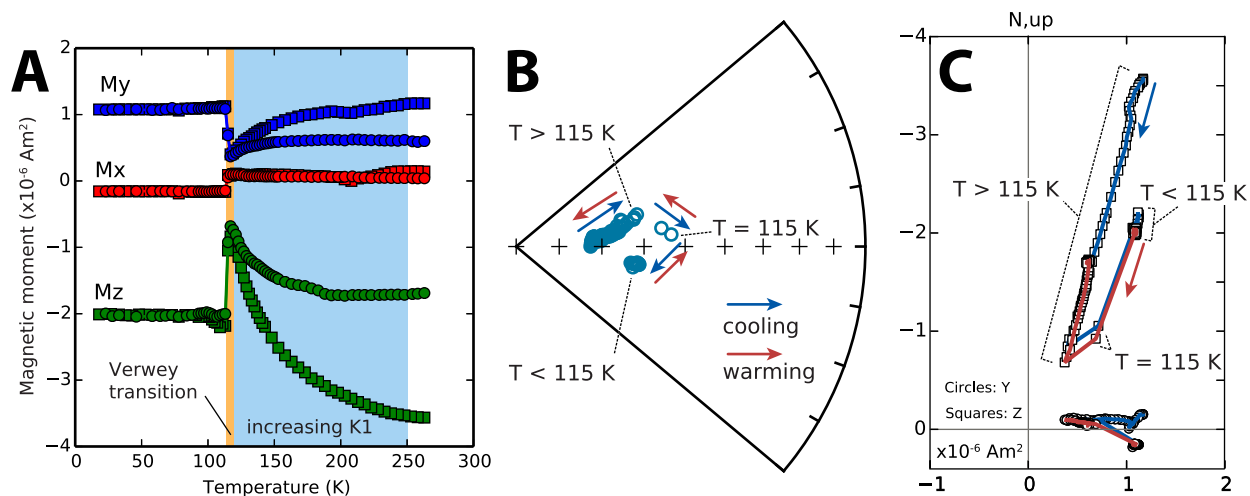
Figure 4b shows the directional results of the low-temperature cycling of the three separate IRM experiments, while Figure 4c shows the magnitude of total magnetization as a function of temperature (as opposed to the  $x$ ,  $y$ , and  $z$  components). The alignment of the IRM remanence with the intended magnetocrystalline axes of the magnetite is broadly similar to the experimental scheme. There are no dramatic changes in remanence direction during low-temperature cycling regardless of the orientation of the applied IRM. During cooling from room temperature the remanence of all three IRMs decreases toward the isotropic point ( $T_K = 130$  K) and Verwey transition ( $T_V = 110$ – $120$  K), reaching a minimum at  $\sim 118$  K (Figure 4c). The



**Figure 4.** Low-temperature thermal cycling of a specimen comprised of a single magnetite crystal. (a) Schematic of the specimen depicting the orientation of the crystal axes along with the directions of the isothermal pulse magnetizations applied during the three experiments. (b) Equal-area plot of the directional data from each experiment. (c) Full vector magnetization intensity for each of the three experiments during cooling (squares) and warming (circles).

extent of this remanence drop is slightly directionally dependent: the percent loss in remanence from 260 K to the Verwey transition minima for the IRMs imparted along  $[-111]$ ,  $[100]$ , and  $[0-1-1]$  are 68%, 75%, and 72%, respectively. Thus, the remanence imparted along the cubic magnetocrystalline easy axis was marginally more resistant to low-temperature demagnetization than those imparted along the intermediate or hard axes. On continued cooling through  $T_v$ , the remanence increases abruptly and significantly (Figure 4c), unlike the behavior commonly observed for polycrystalline MD powder samples in MPMS experiments, but similar to that found in previous single-crystal measurements [e.g., Özdemir and Dunlop, 1999; Smirnov and Tarduno, 2011]. The magnitude of remanence below the Verwey transition is also somewhat anisotropic with the ratio of remanence at 20 K to that at 260 K for each IRM at 53%, 61%, and 56%, respectively. Thus, remanence was more readily inherited by monoclinic magnetite when the original room temperature IRM is applied parallel to its monoclinic magnetocrystalline easy axis direction ( $[001]_m$  or  $[001]_c$ ). A third metric of anisotropy is seen after cycling back to 260 K, where the recovered remanence for the IRMs imparted along  $[-111]$ ,  $[100]$ , and  $[0-1-1]$  is 44%, 48%, and 40%, respectively. In this context, the most efficient remanence recovery occurred when the original IRM was imparted along the cubic magnetocrystalline hard axis. This persistent remanence is likely pinned by crystalline imperfections in the magnetite, such as dislocation networks (as recently imaged by A. Lindquist et al. (Domain wall pinning and dislocations: Investigating magnetite deformed under conditions analogous to nature using transmission electron microscopy, submitted to *Journal of Geophysical Research: Solid Earth*, 2014). The absolute value of the remanence on warming back to room temperature is nearly identical for all three IRM treatments (Figure 4c), suggesting that regardless of the orientation of the initial IRM, after passing twice through  $T_v$  and  $T_K$  the final arrangement of





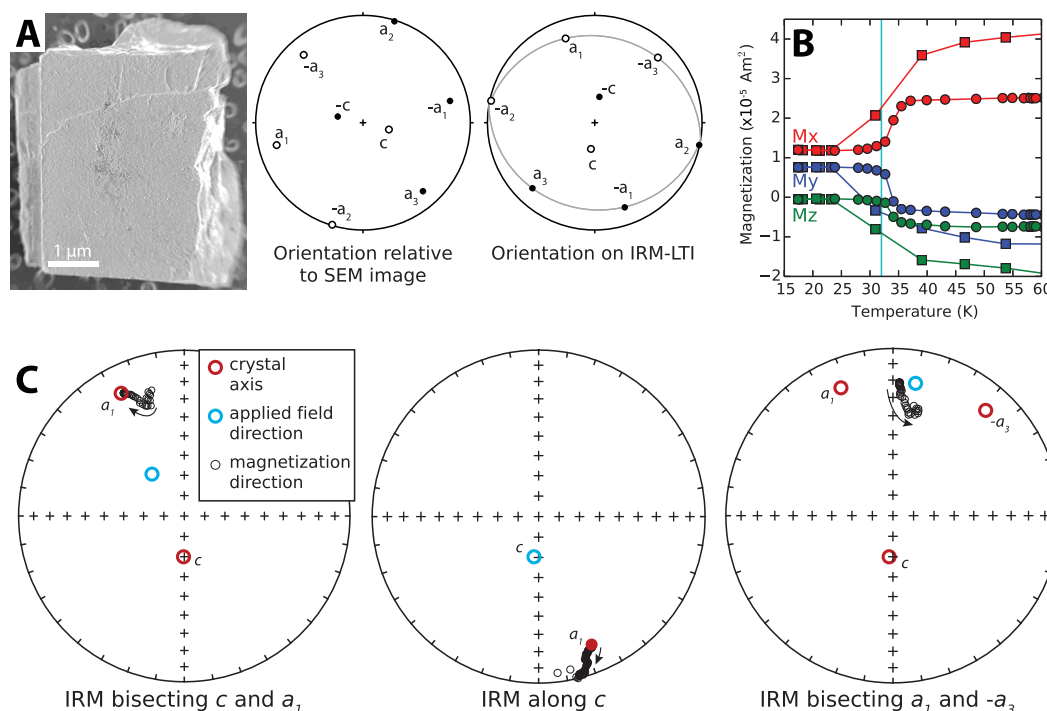
**Figure 5.** Data from the series two experiment (1.2 T IRM applied approximately along the [100] axis) on the single magnetite crystal. (a) Three-axis remanence data for cooling to 20 K and warming back to room temperature using the IRM-LTI. The orange box indicates the interval over which the specimen passes through its Verwey transition, while the blue box indicates the interval of remanence loss as the specimen approaches its isotropic point. (b) Portion of an equal-area plot of the remanence direction during thermal cycling illustrating the clockwise change that occurs as the sample is cooled through the Verwey transition. This directional change is reversible on warming. (c) Vector component diagram where both the change in direction and increase in intensity can be seen across the Verwey transition. The blue line connects data points measured during the cooling portion of the experiment while the red line connects data points for the warming portion of the experiment.

domain walls in the specimen has achieved a common low-energy configuration. These results also show the incomplete nature of low-temperature demagnetization associated with MD-sized grains. However, as is shown in ARM data applied to an assemblage of MD magnetite in a coarse-grained diabase sample below, it is likely that the portion of magnetization that was demagnetized is the lowest coercivity fraction. In this case study, we examine a single, large MD grain and show that a minimum of a third of its remanence is routinely retained after low-temperature cycling. The low-temperature behavior of MD magnetite is further discussed below in section 3.5, and has important implications for paleomagnetic and paleointensity studies that use low-temperature demagnetization steps to minimize the unwanted contributions from MD grains.

The data acquired by the IRM-LTI also allow users to see subtle, but important details that are often lost in one-axis MPMS measurements. For example, the sudden drop in remanence as the magnetite passes through its Verwey transition near 118 K, as well as the more progressive remanence loss as magnetite passes through its isotropic point near 130 K are both well resolved in the three-axis data (Figure 5). Another interesting observation from the three-axis data is the change in the remanence direction in series 2 (where the 1.2 T IRM was applied approximately parallel to the [100] direction) wherein there is subtle clockwise movement on cooling that is reversible on warming. There are two distinct clusters of directions: one cluster is defined by all data collected above 115 K, while the second cluster is comprised of all data collected below 115 K. Measurements collected at 115 K on both cooling and warming are at an intermediate position (Figure 5). The vector component diagram in Figure 5c also shows this behavior along with the concomitant change in intensity before and after the Verwey transition. The crystallographic origins of this directional movement require further study, but one likely influence is the reorganization of ferroelastic twins as the sample is converted from one crystal system to the other [e.g., Kasama *et al.*, 2010].

### 3.4. Single Crystal of Pyrrhotite

A single crystal of pyrrhotite measuring ~2.5 by ~3.5 by ~4.5 mm was used to explore the directional behavior of remanence for multidomain pyrrhotite during low-temperature cycling using the IRM-LTI (Figure 6a). Pyrrhotite often occurs in nature as intergrowths of hexagonal and monoclinic iron sulfide, the latter of which is ferrimagnetic at room temperature. Monoclinic pyrrhotite experiences a dramatic change in its magnetic properties at ~35 K [Besnus and Meyer, 1964] (Figure 6b); the underlying crystallographic and/or magnetocrystalline changes are still debated [Wolfers *et al.*, 2011; Kind *et al.*, 2013]. The details of the mineral's crystal symmetry and internal microstructures are highly complex owing to nonstoichiometry in the Fe<sub>1-x</sub>S system and nonregular ordering (or "incommensurate" ordering) of vacancies within the crystal



**Figure 6.** Experiments conducted on a single crystal of pyrrhotite. (a) Secondary electron SEM micrograph of the sample and crystal orientation information during the progression of an experiment. (b) Three-axis low-temperature remanence data upon cooling (squares) and warming (circles) for the IRM bisecting  $c$  and  $a_1$  experiment. The Besnus transition is well expressed in these data and closely corresponds with the canonical value for the transition (32 K) shown with the cyan vertical line. (c) Equal-area plots showing the directional data (gray circles with arrow indicating direction of change) obtained upon cooling for isothermal remanent magnetizations (IRM) that were applied in three different directions. The direction along which the 1.2 T pulse magnetizations were applied is indicated in each plot with a light blue circle, while relevant crystal axis directions are shown with red circles. Closed/open circles refer to directions in the lower/upper hemisphere.

lattice. Crystallographers frequently take a simplified view of the mineral in referring to its crystal symmetry as “pseudo-hexagonal.” Above the Besnus transition, pyrrhotite exhibits a “hard” magnetic axis along the  $c$  direction and remanence tends to lie within the basal plane defined by the  $a_1$ ,  $a_2$ , and  $a_3$  axes [Martín-Hernández et al., 2008]. A secondary electron micrograph of the single-crystal specimen used in this study is shown in Figure 6a (collected with a JEOL 6500 FEG-SEM operated at 20 kV with a working distance of 22.8 mm). The orientations of the crystal’s pseudo-hexagonal crystallographic axes were measured using electron backscatter diffraction (EBSD) and are shown in Figure 6a. No evidence of monoclinic-hexagonal intergrowths was observed during EBSD analyses, which provides further evidence for the single-crystal nature of this sample. Prior to measurement in the IRM-LTI, the sample was oriented such that the  $a_1$  axis was subparallel to the  $x$  axis of the magnetometer. The remanence of the sample was noted and then the threaded copper sample holder was tightened onto the sample mounting surface to ensure thermal contact, resulting in a counterclockwise rotation of the  $a_1$  axis of 10–30°. IRMs were imparted to the sample in a number of orientations prior to low-temperature cycling in the IRM-LTI, and the intended directions are shown as blue points in the equal-area plots in Figure 6c. As was the case with the single-crystal magnetite measurements, there are likely small angular errors between the actual and intended applied field directions for IRM acquisition.

Figure 6c shows the directions throughout low-temperature demagnetization of three separate 1.2 T IRMs imparted along different crystallographic directions. The leftmost equal-area plot in Figure 6c shows the low-temperature demagnetization of an IRM imparted along an angle bisecting the pyrrhotite  $a_1$  and  $c$  axes. The resulting remanence falls almost directly along the  $a_1$  axis (Figure 6c). This difference in applied field direction and the acquired remanence is likely due to the strong anisotropy given that remanence in pyrrhotite prefers to lie within the basal plane of the pseudo-hexagonal structure. During cooling and warming, the remanence direction remains largely unchanged, but does track slightly along a great circle that includes the  $c$  axis. After passing through the Besnus transition on cooling, the remanence at 20 K is only

10% of the value at 290 K. After warming back to 290 K, the final residual remanence is 18% of the original remanence at 290 K (see supporting information).

The middle equal-area plot in Figure 6c shows an IRM imparted along the  $c$  axis. The resulting remanence falls along the  $-a_1$  direction, again indicating the strong tendency at room temperature for remanence in pyrrhotite to lie within the basal plane. As in the case for the low-temperature cycling the IRM that bisected  $c$  and  $a_1$ , the remanence direction remained relatively stable during low-temperature cycling. However, after passing through the Besnus transition on cooling, the remanence of this IRM at 20 K is 25% of the value at 290 K. After warming back to 290 K, the final residual remanence is 17% of the original remanence at 290 K (Figure 6b). Thus, an IRM applied along the pyrrhotite  $c$  axis allows for a greater percentage of its remanence to be retained after cooling through the Besnus transition.

The rightmost equal-area plot in Figure 6c shows an IRM imparted along an angle that bisects the  $a_1$  and  $-a_3$  directions. In this instance, the resulting remanence direction appears to be oriented roughly parallel to the applied field direction, suggesting that pyrrhotite remanence is not constrained to lie solely along the  $a_i$  axes, but may be oriented subparallel to the applied field within the basal plane. After passing through the Besnus transition on cooling, the remanence at 20 K is 18% of the value at 290 K. After warming back to 290 K, the final residual remanence is 22% of the original remanence at 290 K. Thus, a higher percentage of remanence seems to be retained after cycling through the Besnus transition if an IRM is applied along a direction within the basal plane.

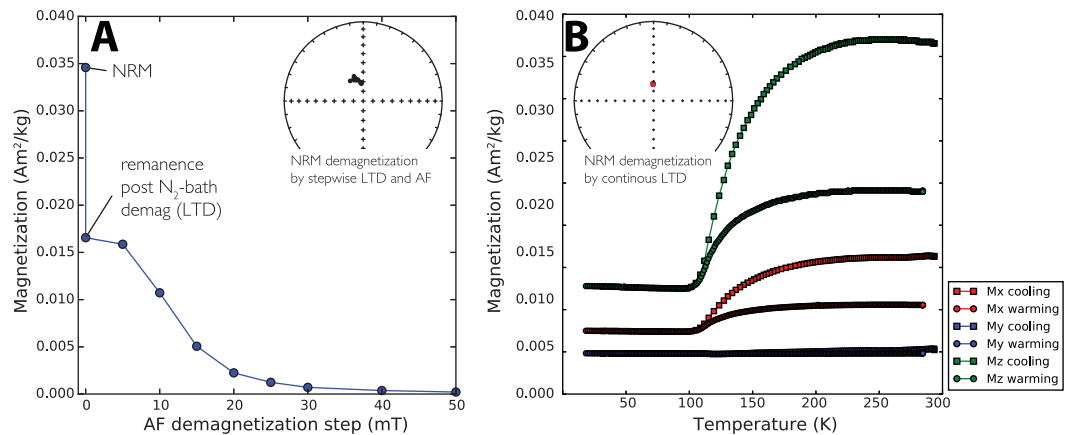
It is worth noting that no evidence of a self-reversal was observed in our IRM experiments on pyrrhotite. This phenomenon has been observed to occur in large single crystals of pyrrhotite that are either internally twinned [e.g., Zapletal, 1992] or partially oxidized [e.g., Bina and Daly, 1994], whereas the crystal the pyrrhotite single crystal in this study is untwinned and unoxidized.

### 3.5. Samples From the Interior of Umkondo Large Igneous Province Sills

Many rocks carry complex multicomponent natural remanent magnetizations, with different generations of magnetization acquired at different times by separate mechanisms and carried in varying proportions by different populations of magnetic grains. For such heterogeneous magnetizations, significant directional changes may accompany passage through transitions and isotropic points as different components are affected according to the mineralogy, defect structures, and grain sizes of their carriers. Just as continuous high-temperature thermal demagnetization can provide crucial observations for correctly interpreting a sample's remanence [Wack and Matzka, 2007; Coe et al., 2014], full-vector measurements during low-temperature cycling of such samples may also provide a new means of disentangling a rock's complex magnetization history.

Bulk rock samples containing large populations of randomly oriented magnetic crystals carrying a univectorial remanence (i.e., a magnetization acquired through a single process, in a relatively constant field orientation), are not expected to show large directional changes while passing through magnetic transitions. In such a scenario, the directional changes associated with the individual crystals should average out, and only a change in the magnitude of the remanence should be observed. This is the case for specimen PW15-4d, a coarse-grained diabase sill from the  $\sim 1.1$  Ga Umkondo Province of Botswana [Hanson et al., 2004]. Room temperature alternating-field demagnetization experiments show that the NRM decays in a univectorial manner (Figure 7a). Traditional low-temperature demagnetization experiments involving submersion in a liquid  $N_2$  bath result in a 52% loss in NRM intensity (Figure 7a), suggesting that a significant fraction of the remanence is held by multidomain grains. Low-temperature cycling of the NRM in the IRM-LTI on a sister specimen from the same core shows that the remanence direction remains stationary while the intensity decreases significantly such that it is only 21% of the initial NRM after cooling through the Verwey transition. The magnetization recovers to 52% of the original NRM value upon warming back to room temperature (Figure 7b) resulting in a similar overall remanence loss as the liquid nitrogen bath on the sister specimen.

In order to explore whether low-coercivity magnetite loses its remanence at similar temperatures to higher-coercivity magnetite, this same specimen was given two mutually perpendicular anhysteretic remanent magnetizations (ARMs; Figure 8). First, the specimen was given an ARM along its negative  $z$  axis (200 mT AF with a 100  $\mu$ T DC bias field). Then the specimen was imparted with a second more weakly held ARM along

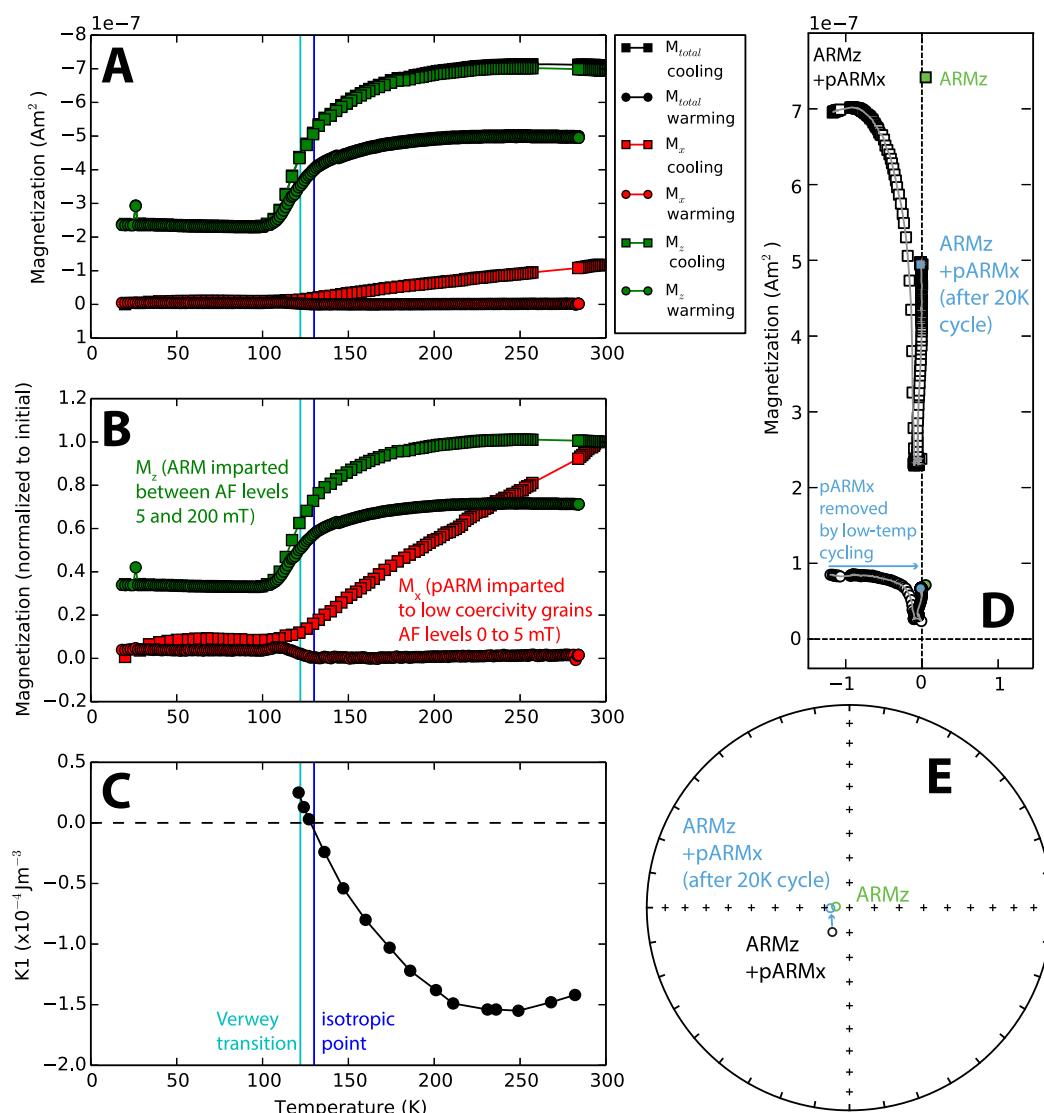


**Figure 7.** Natural remanent magnetization (NRM) data for a coarse-grained diabase sill sample (PW15-4a). (a) Demagnetization data that begin with a low-temperature demagnetization step (a liquid  $N_2$  bath in near-zero field) followed by AF demagnetization. The remanence is comprised of a single directional component as can be seen in the inset equal-area plot. (b) Three-axis low-temperature cycling data of NRM obtained on the IRM-LTI for the natural remanence of a sister specimen from the same sample. The directions are summarized on the inset equal-area plot and are very similar to the directions of the stepwise AF demagnetization data.

its negative  $x$  axis (5 mT AF with a 200  $\mu$ T DC bias field). This second ARM targeted only those grains in the sample with coercivities  $\leq 5$  mT, and hence is a partial ARM (pARM). In contrast, the first ARM applied along the  $z$  axis is held by higher-coercivity grains ranging between 5 and 200 mT (referred to as “ARM $_z$ ” henceforth). On thermal cycling in the IRM-LTI, the remanence associated with ARM $_z$  behaves in a manner similar to what is commonly observed in single-axis MPMS measurements, where remanence begins to decrease at  $\sim 225$  K and then decreases at an accelerating rate as it approaches  $T_V$  (Figure 8a). This pattern of remanence loss parallels the changes in the dominant magnetocrystalline energy term for cubic magnetite ( $K_1$ ) (Figure 8c). The isotropic point for magnetite,  $T_K$ , occurs near 130 K when  $K_1$  changes from negative to positive. Thus, this pattern of remanence loss in magnetite during cooling is actually a two-stage process involving passage through both  $T_K$  and  $T_V$ . During warming, the remanence associated with the ARM $_z$  recovers to 71% of its original value. This low-temperature memory is significant and is typical for single-domain and pseudo-single-domain grains, which are able to preserve their original remanence owing to shape anisotropy and pinned domain walls [Dunlop and Özdemir, 1997]. It is also likely that some of this remanence is held by multi-domain grains with regions of high defect density that produce pinning coercivities higher than 5 mT.

By contrast, the pattern of remanence loss associated with the lower coercivity pARM is significantly different from that of the ARM $_z$  (Figures 8a and 8b). The lower coercivity remanence experiences a linear decrease from room temperature to  $T_V$  that is seemingly unaffected by the rate of change in the  $K_1$  magnetocrystalline energy term (Figure 8c). This style of low-temperature demagnetization behavior is novel and to our knowledge has not been observed in earlier studies of magnetite-bearing rocks, although Muxworthy *et al.* [2003] did observe systematic changes in the shapes of  $M(T)$  cooling curves for pARMs acquired over different coercivity windows, in magnetites of differing grain sizes. Importantly, this behavior would be undetected in a conventional MPMS RT-SIRM measurement. There is no recovery of this remanence during warming back to room temperature (Figures 8a and 8b), suggesting that this remanence was held by multi-domain grains whose domain wall configurations were dramatically rearranged during cycling through both  $T_K$  and  $T_V$ . The final remanence direction following low-temperature cycling of the ARM $_z$  + pARM $_x$  is very similar to the ARM $_z$  direction prior to imparting pARM $_x$ . This result demonstrates the effectiveness of low-temperature demagnetization for the removal of remanence held by low-coercivity grains.

Some of the specimens from the Umkondo province sills show significant directional changes after thermal cycling in liquid  $N_2$ . Figure 9 shows one such specimen, PW10-7a, whose NRM direction shifted by  $64^\circ$  after the low-temperature demagnetization (LTD) step. A conventional MPMS RT-SIRM measurement (see supporting information) shows that 67% of the room temperature remanence is lost through low-temperature demagnetization. However, the IRM-LTI provides a more complete picture, capturing this same broad behavior, but also documenting the progressive directional change throughout the experiment (Figure 9). After passing through  $T_V$ , the remanence direction remains relatively constant during cooling to 20 K and subsequent rewarming to  $T_V$ . However, during warming from  $T_V$  to room temperature the specimen's



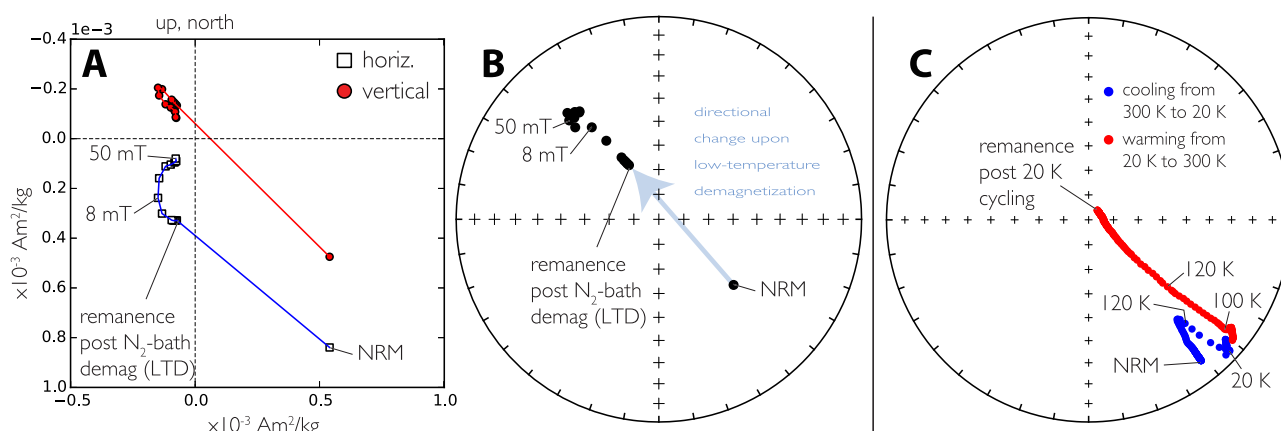
**Figure 8.** Low-temperature behavior of an ARM applied to the z axis and a pARM applied to the x axis of a specimen of an Umkondo dike sill (PW15-4d). (a) The magnetization of the z and x components as well as the total magnetization upon cooling (note that the y axis of the plot is inverted for comparison between the components and the total magnetization). The z and x components are normalized to their initial magnetizations in Figure 8b to highlight the differing behavior between the x axis magnetization (pARM imparted to low-coercivity grains) and the z axis magnetization (ARM on the rest of the coercivity spectrum). (c) The magnetocrystalline anisotropy constant  $K_1$  as a function of temperature (data of Bickford et al. [1957] as presented in Muxworthy and McClelland [2000]). (d) The vector-component diagram and (e) equal-area plot show the directional change associated with low-temperature cycling and illustrate the return to the direction of the z axis ARM with the x axis pARM demagnetized.

remanence direction shifts dramatically toward expected Umkondo orientations as the low-temperature memory held by SD and PSD grains is preferentially recovered over magnetization held by MD grains that is not recovered. As a result, the overprint, secondary to the primary thermal remanent magnetization is largely, although not entirely, removed. These three-axis data demonstrate the effectiveness of low-temperature cycling as a tool for demagnetization of remanence held by MD grains. Furthermore, these data give a glimpse into the “black box” of progressive remanence change that occurs during routine liquid  $\text{N}_2$  demagnetization of paleomagnetic samples.

#### 4. Future Directions

This study presents several examples demonstrating the utility of low-temperature remanence cycling using the IRM-LTI, but there are a multitude of additional research directions that can be explored using this





**Figure 9.** Natural remanent magnetization (NRM) data for a coarse-grained diabase sill sample (PW10-7). (a) Demagnetization data that begin with a low-temperature demagnetization step (a liquid  $N_2$  bath) followed by AF demagnetization. The vector component diagram in Figure 9a and (b) equal-area plot show significant loss of remanence and directional change that accompanied the liquid  $N_2$  treatment. (c) Low-temperature cycling of the NRM of a sister specimen on the IRM-LTI-documented directional change throughout the experiment with movement in the direction of overprint removal both during cooling to the Verwey transition and upon warming back across the Verwey transition.

instrument. From an experimental perspective, one of the more exciting applications for the IRM-LTI is the potential for in-line treatments at low temperatures. Examples include alternating-field demagnetization and acquisition of anhysteretic remanence magnetization (ARM) or IRM. Such treatments would allow researchers to explore changes in the coercivity distributions of samples above and below important mineral transitions, such as the *Morin*, *Verwey*, and *Besnus* transitions. For nanoparticle populations, the joint distribution of grain size and microcoercivity  $f(V, H_{ko})$  can be obtained by AF demagnetization of weak-field TRMs at a set of low temperatures [cf. *Dunlop and West, 1969*]. To enable such new experimental capabilities, the construction of in-line coils is currently under way at the IRM. Additionally, as alluded to in section 3.2, the IRM-LTI also has the potential to run three-axis Lowrie style thermal cycling experiments that can enable users to unmix the low-temperature remanence held by populations of different phases and grain sizes with greater confidence than possible with a single-axis instruments. This approach will provide information analogous to that obtained in traditional Lowrie tests, but without the potential for mineral alteration typically associated with thermal demagnetization experiments above room temperature.

The experimental flexibility of the IRM-LTI may make it possible to address many of the questions regarding fundamental magnetic behavior in important minerals such as magnetite, maghemite, titanomagnetite, hematite, greigite, and pyrrhotite. The IRM-LTI allows researchers to observe full-vector remanence changes as minerals pass through isotropic points, ordering temperatures, and blocking and unblocking temperatures. Processes such as TRM acquisition can be explored by measuring the remanence acquired by populations of superparamagnetic grains. Similarly, the effects of cooling rate on the intensity of remanence can be directly observed by modifying the rate at which specimens are cooled. By preparing specimens with varying degrees of mineral alignment it may also be possible to quantify the anisotropy of remanence. This is by no means an exhaustive exposition of the possible rock magnetic applications, but we hope that it illustrates the research potential of the instrument.

The IRM-LTI can also provide complementary information to planetary scientists trying to isolate the primary magnetization of meteorites in order to learn more about the conditions that existed during the early history of the solar system. The assemblage of magnetic minerals found within meteorite samples is frequently far more complex than that found in terrestrial samples. Alloys of Fe, Ni, and S are common, as are varying concentrations of magnetite and titanomagnetite. The effects of shock on remanent magnetization further compound the difficulties of interpreting a meteorites complex thermomagnetic history. One process that has not received as much attention is the effect of low-temperature thermal cycling on the remanent magnetism of the constituent magnetic minerals. The orbits of most asteroids and comets are eccentric, which guarantees that these objects will experience heating and cooling cycles during the course of their orbit around the Sun. Another process that may affect the magnetization of meteorites is inverse thermoremanent magnetization, which is acquired as magnetite is warmed through the Verwey transition or isotropic point [Dunlop, 2006]. This type of remanence could be acquired in Earth's field as meteorites warm to Earth



surface temperatures and could contaminate paleointensity measurements. The IRM-LTI's low-temperature, three-component system allows researchers to design experiments to observe remanence changes that could result from thermal perturbations experienced by meteorites as they pass by the Sun (at solar perigee) or as they enter Earth's atmosphere.

These applications form the basis for decades of potential future research; much more than could ever be accomplished by the authors of this study. The instrument was designed with a philosophy of making it affordable to other research groups and to be broadly compatible with horizontal-loading 2G Enterprises cryogenic magnetometers that are used in the majority of paleomagnetic laboratories. The IRM-LTI is currently housed at the Institute for Rock Magnetism at the University of Minnesota, where it is available to the international scientific community through the lab's NSF Visiting Fellowship program (<http://www.irm.umn.edu/IRM/categories.html>). It is hoped that this study will inspire colleagues to visit the IRM in order to use the instrument to advance their own research programs.

### Acknowledgments

Bill Goree's career of developing and commercializing superconducting rock magnetometers for paleo and rock magnetism continues to leave a substantial impact on our discipline. This project builds on his vision of a low-temperature insert that just before his passing in 2007 he envisioned would be best implemented with a liquid-free cryocooler system. John Tarduno encouraged the development of this type of insert and provided valuable input based on his lab's experience with their flow-through system. Richard Hanson played an important role in the field expedition during which samples were collected from Umkondo Sills. Functions from the open source PmagPy project (<https://github.com/Itauxe/PmagPy>) were used extensively for data plotting, and we are grateful to the project's lead developer Lisa Tauxe. The manuscript benefited from reviews by Mark Dekkers, Aleksey Smirnov, and one anonymous researcher. This research was funded by EAR-0929807 to J.M.F., EAR-1045635 to N.L.S.-H., and by support of the IRM by the NSF/EAR Instruments and Facilities program. This is IRM contribution 1409.

### References

- Besnus, M. J., and A. J. Meyer (1964), Nouvelles données expérimentales sur le magnétisme de la pyrrhotine naturelle, in *Proceedings of International Conference on Magnetism*, vol. 20, pp. 507–511, Inst. of Phys. and the Phys. Soc., London, U. K.
- Bickford, L. R., J. M. Brownlow, and R. F. Penoyer (1957), Magnetocrystalline anisotropy in cobalt-substituted magnetite single crystals, *IEE Proc., Part B*, 104(5), 238–244, doi:10.1049/pi-b-1.1957.0038.
- Bina, M., and L. Daly (1994), Mineralogical change and self-reversed magnetizations in pyrrhotite resulting from partial oxidation; geophysical implications, *Phys. Earth Planet. Inter.*, 85(1–2), 83–99, doi:10.1016/0031-9201(94)90009-4.
- Coe, R. S., N. A. Jarboe, M. Le Goff, and N. Petersen (2014), Demise of the rapid-field-change hypothesis at Steens Mountain: The crucial role of continuous thermal demagnetization, *Earth Planet. Sci. Lett.*, 400, 302–312, doi:10.1016/j.epsl.2014.05.036.
- Dunlop, D. and Ö. Özdemir (1997), *Rock Magnetism: Fundamentals and Frontiers*, Cambridge Univ. Press, Cambridge, U. K.
- Dunlop, D. J. (2003), Stepwise and continuous low-temperature demagnetization, *Geophys. Res. Lett.*, 30(11), 1582, doi:10.1029/2003GL017268.
- Dunlop, D. J. (2006), Inverse thermoremanent magnetization, *J. Geophys. Res.*, 111, B12S02, doi:10.1029/2006JB004572.
- Dunlop, D. J., and G. F. West (1969), An experimental evaluation of single domain theories, *Rev. Geophys.*, 7(4), 709–757, doi:10.1029/RG007i004p00709.
- Frederichs, T., T. von Dobeneck, U. Bleil, and M. J. Dekkers (2003), Towards the identification of siderite, rhodochrosite, and vivianite in sediments by their low-temperature magnetic properties, *Phys. Chem. Earth*, 28(16–19), 669–679, doi:10.1016/S1474-7065(03)00121-9.
- Hanson, R., J. Crowley, S. Bowring, J. Ramezani, W. Gose, I. Dalziel, J. Pancake, E. Seidel, T. Blenkinsop, and J. Mukwakwami (2004), Coeval large-scale magmatism in the Kalahari and Laurentian Cratons during Rodinia assembly, *Science*, 304, 1126–1129, doi:10.1126/science.1096329.
- Kasama, T., N. S. Church, J. M. Feinberg, R. E. Dunin-Borkowski, and R. J. Harrison (2010), Direct observation of ferrimagnetic/ferroelastic domain interactions in magnetite below the Verwey transition, *Earth Planet. Sci. Lett.*, 297(1–2), 10–17, doi:10.1016/j.epsl.2010.05.004.
- Kind, J., I. García-Rubio, M. Charilaou, N. R. Nowaczyk, J. F. Löffler, and A. U. Gehring (2013), Domain-wall dynamics in 4C pyrrhotite at low temperature, *Geophys. J. Int.*, 195(1), 192–199, doi:10.1093/gji/ggt262.
- Lowrie, W. (1990), Identification of ferromagnetic minerals in a rock by coercivity and unblocking temperature properties, *Geophys. Res. Lett.*, 17(2), 159–162.
- Martin-Hernández, F., M. J. Dekkers, I. M. A. Boinaars-Silken, and J. C. Maan (2008), Magnetic anisotropy behaviour of pyrrhotite as determined by low- and high-field experiments, *Geophys. J. Int.*, 174(1), 42–54, doi:10.1111/j.1365-246X.2008.03793.x.
- Morin, F. J. (1950), Magnetic susceptibility of  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub> and  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub> with added titanium, *Phys. Rev.*, 78, 819–820.
- Muxworthy, A. R., and E. McClelland (2000), Review of the low-temperature magnetic properties of magnetite from a rock magnetic perspective, *Geophys. J. Int.*, 140, 101–114.
- Muxworthy, A. R., D. J. Dunlop, and Ö. Özdemir (2003), Low-temperature cycling of isothermal and anhysteretic remanence: Microcoercivity and magnetic memory, *Earth Planet. Sci. Lett.*, 205(3–4), 173–184, doi:10.1016/S0012-821X(02)01039-7.
- Özdemir, Ö., and D. J. Dunlop (1999), Low-temperature properties of a single crystal of magnetite oriented along principal magnetic axes, *Earth Planet. Sci. Lett.*, 165(2), 229–239, doi:10.1016/S0012-821X(98)00269-6.
- Schmidt, P. W. (1993), Palaeomagnetic cleaning strategies, *Phys. Earth Planet. Inter.*, 76(1–2), 169–178, doi:10.1016/0031-9201(93)90066-I.
- Smirnov, A. V., and J. A. Tarduno (2011), Development of a low-temperature insert for the measurement of remanent magnetization direction using superconducting quantum interference device rock magnetometers, *Geochem. Geophys. Geosyst.*, 12, Q04Z23, doi:10.1029/2011GC003517.
- Verwey, E. J. W. (1939), Electronic conduction of magnetite (Fe<sub>3</sub>O<sub>4</sub>) and its transition point at low temperatures, *Nature*, 144, 327–328.
- Wack, M., and J. Matzka (2007), A new type of a three-component spinner magnetometer to measure the remanence of rocks at elevated temperature, *Earth Planets Space*, 59, 853–862.
- Wolters, P., G. Fillion, B. Ouladdiaf, R. Ballou, and P. Rochette (2011), The pyrrhotite 32K magnetic transition, *Solid State Phenomena*, 170, 174–179, doi:10.4028/www.scientific.net/SSP.170.174.
- Worm, H.-U., and M. Jackson (1999), The superparamagnetism of Yucca Mountain Tuff, *J. Geophys. Res.*, 104(B11), 25,415–25,425.
- Yamamoto, Y., H. Tsunakawa, and H. Shibuya (2003), Palaeointensity study of the Hawaiian 1960 lava: Implications for possible causes of erroneously high intensities, *Geophys. J. Int.*, 153(1), 263–276.
- Zapletal, K. (1992), Self-reversal of isothermal remanent magnetization in a pyrrhotite (Fe<sub>7</sub>S<sub>8</sub>) crystal, *Phys. Earth Planet. Inter.*, 70(3–4), 302–311, doi:10.1016/0031-9201(92)90196-3.